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RADIO WAVE PROPAGATION
IN AN INHOMOGENEOUS ATMOSPHERE

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U. S. DEPARTMENT OF COMMERCE
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ABSTRACT

A self-contained treatment of the theory of radio waves in an inhomogeneous atmosphere is given. The refractive index is assumed to vary with height above the earth in a monotonic fashion. Variation according to an exponential law is used for illustration of general principles. For this case, rigorous series formulas are developed for the distance to the horizon for an elevated point in the atmosphere.



ON RADIO WAVE PROPAGATION IN AN INHOMOGENEOUS ATMOSPHERE

1. Introduction and Formulation

It is the purpose of this paper to review, and extend, work of Watson¹, Eckersley and Millington², Bremmer³, Vvedensky⁴ and Fock⁵ on propagation of radio waves in an inhomogeneous atmosphere. In particular an attempt will be made to present a unified treatment which includes the analyses of the above authors and shows that "the distance to the horizon" is a valid concept in diffraction theory even when the atmosphere is not homogeneous. The earth is to be represented by a homogeneous sphere of radius a and the source is to be a vertical electric dipole which for the moment, is located near the earth's surface. The usual spherical coordinates (r, θ, ϕ) are chosen with the source oriented along the polar axis and the surface of the earth is defined by $r = a$. The dielectric constant relative to free space is denoted by K and is understood to be a function of r and decreases monotonically with increasing r .

The resultant magnetic field has only a single component, H_ϕ , because of the assumed polar symmetry in the problem. As a further consequence of the symmetry, the electric field component E_ϕ is zero. It is convenient to express the remaining field components in terms of single scalar function U which is actually the radial component of a Hertz vector. For a time factor $\exp(i\omega t)$, this leads to

$$H_\phi = ik_0 K \frac{\partial U}{\partial \theta} , \quad (1)$$

$$E_r = \frac{120\pi}{r \sin \theta} \frac{\partial}{\partial \theta} \left(\sin \theta \frac{\partial U}{\partial \theta} \right), \text{ and} \quad (2)$$

$$E_\theta = \frac{120\pi}{K r} \frac{\partial}{\partial r} \left(K r \frac{\partial U}{\partial \theta} \right) \quad (3)$$

where $k_0 = 2\pi/\text{wavelength in vacuo}$.

These field components satisfy Maxwell's equations if the function U satisfies

$$\frac{\partial}{\partial r} \left[\frac{1}{K} \frac{\partial}{\partial r} (K r U) \right] + \frac{1}{r \sin \theta} \frac{\partial}{\partial \theta} \left[\sin \theta \frac{\partial U}{\partial \theta} \right] + k_0^2 K r U = 0 \quad (4)$$

Following the work of Fock⁵, it now proves to be convenient to introduce an auxiliary quantity U' which is defined by

$$U' = e^{-ika\theta} K r \sqrt{\sin \theta} U \quad (5)$$

and consequently the equation for U satisfies

$$\begin{aligned} \frac{\partial^2 U'}{\partial r^2} - \frac{2ik}{a} \frac{\partial U'}{\partial \theta} + k^2 \left(\frac{K}{K_0} - \frac{a^2}{r^2} \right) U' \\ = \frac{K'}{K} \frac{\partial U'}{\partial r} - \frac{1}{r^2} \left[\frac{\partial^2 U'}{\partial \theta^2} + \left(\frac{1}{4} + \frac{1}{r \sin^2 \theta} \right) U' \right] \end{aligned} \quad (6)$$

where $K' = dK/dh$, K_0 is the value of K at $r = a$, and $k = k_0 \sqrt{K_0}$ is the effective wave number at the surface of the earth.

Since $E_\theta \cong -ZH_\phi$ at $r = a$ it follows that

$$\frac{\partial U'}{\partial r} = ik (K_0)^{\frac{1}{2}} Z U' \text{ at } r = a \quad (7)$$

where Z is the effective value of the surface impedance⁶⁻⁸.

The right-hand side of equation (6) can be replaced by zero. The justification for doing this is based on the fact that

$$\partial U' / \partial r = \text{Order of } (kU'/M)$$

and

$$\partial U' / \partial \theta = \text{Order of } (kaU'/M^2)$$

where $M = (ka/2)^{1/3}$ is a large quantity. It is also possible to simplify equation (6) by noting that

$$(1 - a^2/r^2) \cong 2h/a .$$

Consequently U' satisfies

$$\frac{\partial^2 U'}{\partial h^2} - 2ik \frac{\partial U'}{\partial s} + k^2 \left(\frac{K - K_0}{K_0} + \frac{2h}{a} \right) U' = 0 \quad (8)$$

where $s = a\theta$ is the great-circle distance measured along the earth's surface. The boundary condition at the surface of the earth then becomes

$$\partial U' / \partial h = +ik (K_0)^{\frac{1}{2}} Z U' \quad (\text{at } h = 0) \quad (9)$$

The function U' , which must satisfy equations (8) and (9), should also have the proper singularity at the source. That is it should behave as

$$U = \frac{e^{-ikR}}{R} (1 + R_v) \quad (10)$$

near the source, where R_v is a Fresnel reflection coefficient of the ground for vertically polarized waves. For $h \ll s$, and assuming that our approximate boundary condition is valid, $R \cong s + (h^2/2s)$

and thus

$$R_v \cong \frac{h/s - Z}{h/s + Z} \quad (11)$$

Consequently, the required behavior of U' near the origin is

$$U' \cong K_0 (a/s)^{\frac{1}{2}} (2h/s) (h/s + Z)^{-1} e^{-i \frac{kh^2}{2s}} \quad (12)$$

2. Equivalent Earth Radius Concept

The equivalent earth radius a_e will now be defined in the following manner

$$\frac{1}{a_e} = \frac{1}{a} + \frac{K'_0}{2K_0} \quad \text{where} \quad K'_0 = \left(\frac{\partial K}{\partial h} \right)_{h=0} \quad (13)$$

This equation is usually justified on the basis of geometrical optics. It can also be demonstrated in a more rigorous fashion by rewriting equation (8) in the form

$$\frac{\partial^2 U'}{\partial h^2} - 2ik \frac{\partial U'}{\partial s} + \frac{2k^2}{a_e} h (1 + g) U' = 0 \quad (14)$$

where

$$g = \frac{a_e}{2K_o} \left(\frac{K - K_o}{h} - K'_o \right) .$$

Now if the atmosphere is linear then by definition $K'_o = (K - K_o)/h$ and $y = 0$ and consequently

$$\frac{\partial^2 U'}{\partial h^2} - 2ik \frac{\partial U'}{\partial s} + \frac{2k^2}{a_e} h U' = 0 . \quad (15)$$

But, on the other hand, if $K = K_o$ for all h then from equation (8) or (14),

$$\frac{\partial^2 U'}{\partial h^2} - 2ik \frac{\partial U'}{\partial s} + \frac{2k^2}{a} h U' = 0 . \quad (16)$$

This strongly suggests that propagation in a linear atmosphere of radius a is equivalent to propagation over an earth of modified radius a_e with a homogeneous atmosphere. It is therefore concluded that the effective earth radius is indeed a valid concept. This conclusion has also been reached by many other authors, including Bremmer,³ Fock⁵ and Miller⁹.

3. Extension to Non-Linear Atmosphere

In many practical cases of interest, however, K is not a linear function of height, particularly at great heights. Then the function g is non-zero and is itself, in general, a function of height. To further analyze this case, the relevant equations will be put in dimensionless form. The following parameters are first introduced

$$x = \left(\frac{ka_e}{2} \right)^{\frac{1}{3}} \quad \frac{s}{a_e} = \left(\frac{ka_e}{2} \right)^{\frac{1}{3}} \theta ,$$

$$y = \left(\frac{2}{ka_e} \right)^{\frac{1}{3}} kh ,$$

$$q = -i \left(\frac{ka_e}{2} \right)^{\frac{1}{3}} (K_o)^{\frac{1}{2}} (Z/120\pi), \text{ and}$$

$$W = \frac{1}{K_o} \left(\frac{2}{ka_e} \right)^{1/6} U' .$$

The equation for W is

$$\frac{\partial^2 W}{\partial y^2} - i \frac{\partial W}{\partial x} + y (1 + g) W = 0 \quad (17)$$

and the boundary condition at the earth's surface now becomes

$$\frac{\partial W}{\partial y} + q W = 0 \text{ at } y = 0 . \quad (18)$$

Furthermore, as x tends to zero, W should behave as $(2/\sqrt{x}) \cdot \exp(-iy^2/4x)$.

It is now desirable to first consider the solution of equation (17) when y can be neglected compared to unity. This corresponds to an atmosphere which is linear or near linear. The relevant equation is now

$$\frac{\partial^2 W}{\partial y^2} - i \frac{\partial W}{\partial x} + y W = 0 . \quad (19)$$

This can be solved by the method of separation of variables. That is, assuming

$$W = X(x) Y(y)$$

one finds

$$\frac{1}{Y(y)} \frac{d^2 Y(y)}{dy^2} = i \frac{1}{X(x)} \frac{d X(x)}{dx} = t \quad (20)$$

where t the separation parameter. Thus

$$\frac{dX}{dx} + it X = 0 \quad (21)$$

and

$$\frac{d^2 Y}{dt^2} + (y - t) Y = 0 . \quad (22)$$

Solutions of the first equation are

$$X(x) = e^{\pm itx} \quad (23)$$

and the second equation has solutions of the form

$$Y(y) = W(t - y) \quad (24)$$

where $W(t)$ satisfies Airy's equation

$$\frac{d^2 W(t)}{dt^2} - t W(t) = 0 . \quad (25)$$

The appropriate solution for the x dependence is chosen such that outgoing waves are only retained. Therefore

$$X(x) = e^{-itx} \quad (26)$$

Furthermore, the radial part of the solution (i. e. y dependence) is chosen so that the phase of $Y(y)$ is a decreasing function with height (i. e. phase lag increases with height). At this stage, it should be noted that Airy's equation is closely related to Bessel's equation of order $1/3$. In fact

$$W(t) = \text{const.} \times t^{\frac{1}{2}} Z_{1/3} [(2/3) (-t)^{2/3}] \quad (27)$$

where $Z_{1/3}$ is any cylindrical Bessel function or linear combinations thereof. In order to satisfy our requirement that the phase lag be an increasing function with height, the Hankel function $H_{1/3}^{(2)}$ of the second must be used. A particular solution which is suitable from the above standpoint is

$$W(t) = W_1(t) = e^{-i2\pi/3} (-\pi t/3)^{\frac{1}{2}} H_{1/3}^{(2)} [(2/3) (-t)^{3/2}] \quad (28)$$

where the constant is chosen to be consistent with the work of Fock and Leontovich. In fact, the latter authors represent $W_1(t)$ as an integral of the Airy type

$$W_1(t) = \frac{1}{\sqrt{\pi}} \int_{\infty}^{\infty} e^{i2\pi/3} e^{tz - z^{3/3}} dz \quad (29)$$

where the contour is a straight line from infinity along the straight $\arg. t = 2\pi/3$ to the origin and then out to ∞ along the real axis. (i. e. $\arg. t = 0$).

To satisfy the boundary condition at the earth's surface, one must choose

$$[W_1'(t-y) - q W_1(t-y)]_{y=0} = 0 \quad (30)$$

A numerical study of this equation indicates that the roots $t = t_s$ ($s = 0, 1, 2, \dots$) are in the fourth quadrant of the t plane. The solution is then made up a linear combination of terms of the type $e^{-ixt_s} W_1(t_s - y)$.

It can be seen immediately that an equivalent representation is

$$W_1 = \gamma \int_C \frac{e^{-ixt} W_1(t-y)}{W_1'(t) - q W_1(t)} dt \quad (31)$$

where C is some contour which encloses the poles of the integrand (i. e. at $t = t_g$). The factor γ , assumed constant, is to be chosen such that W_1 has the proper singularity as x approaches zero. It is to be expected that behavior of W_1 for small x is to be determined by the behavior of the integrand for large values of t because of the factor e^{-ixt} . But, when t is very large, it follows from the known asymptotic behavior of the Hankel functions, that

$$\frac{W_1(t-y)}{W_1'(t) - q W_1(t)} \cong \pm (t)^{-\frac{1}{2}} \exp[-y(t)^{\frac{1}{2}}] \quad (32)$$

where the upper (+) sign is to be employed for $2\pi/3 \geq \arg t > -\pi/3$ and the lower (-) sign for $-\pi/3 > \arg t \geq 4\pi/3$. The contour C is now drawn such that one branch is along the negative imaginary axis ($-i\infty$ to 0) and the other along the real axis (0 to ∞). For small values of x , then

$$W \sim \gamma \int_0^{-i\infty} e^{-ixt} \frac{e^{y\sqrt{t}}}{\sqrt{t}} dt + \gamma \int_0^{\infty} e^{-ixt} \frac{e^{-y\sqrt{t}}}{\sqrt{t}} dt \quad (33)$$

These integrals can be transformed to the form

$$\int_0^{\infty} e^{-\beta^2} d\beta = \sqrt{\pi}/2 \quad (34)$$

and thus

$$W \sim 2\gamma \left(\frac{\pi}{ix}\right)^{\frac{1}{2}} e^{-i\frac{y^2}{4x}} \quad (35)$$

and therefore, if $\gamma = (i/\pi)^{\frac{1}{2}}$

$$W \cong 2 x^{-\frac{1}{2}} e^{-i \frac{y^2}{4x}} \quad (36)$$

as x approaches zero.

The required solution can thus be written

$$W = \left(\frac{i}{\pi}\right)^{\frac{1}{2}} \int_C \frac{e^{-ixt} W_1(t-y)}{W_1(t) - q W_1(t)} dt \quad (37)$$

In this particular case the solution for U becomes

$$U = \frac{e^{-ika\theta}}{a (\theta \sin \theta)^{\frac{1}{2}}} V \cong \frac{e^{-ika\theta}}{a\theta} V \quad (38)$$

where $V = x^{\frac{1}{2}} W$

$$V = \left(\frac{ix}{\pi}\right)^{\frac{1}{2}} \int_C \frac{e^{-ixt} W_1(t-y)}{W_1(t) - q W_1(t)} dt \quad (39)$$

Attention is now returned to the case of an inhomogeneous atmosphere wherein $g(y)$ is no longer negligible compared to unity. The more general equation for W must then be retained

$$\frac{\partial^2 W}{\partial y^2} - i \frac{\partial W}{\partial x} + y [1 + g(y)] W = 0. \quad (40)$$

Proceeding in an analogous fashion as in the linear case, it is readily found that solutions are of the form

$$W = e^{-ixt} F(y, t) \quad (41)$$

where $F(y, t)$ must satisfy

$$\frac{d^2 F(y, t)}{dy^2} + [y - t + y g(y)] F(y, t) = 0. \quad (42)$$

The solution is then expressed as an integral

$$W_1 = \sqrt{\frac{i}{\pi}} \int_C \frac{e^{-ixt} F(y, t)}{\left[\frac{\partial F}{\partial y} + qF \right]_{y=0}} dt \quad (43)$$

which is constructed so the poles of the integrand are the roots,

$$t = t_g, \text{ of the equation } \left[\frac{\partial F(y, t)}{\partial y} + q F(y, t) \right]_{y=0} = 0 \quad (44)$$

and the constant term outside the integral is chosen so that it reduces to the corresponding expression for W_1 in the linear atmosphere (i. e. $g = 0$). The corresponding (formal) solution for the function U is thus written

$$U = \frac{e^{-ika\theta}}{a (\theta \sin \theta)^{\frac{1}{2}}} V \quad (45)$$

where

$$V = \left(\frac{ix}{\pi} \right)^{\frac{1}{2}} \int_C \frac{e^{-ixt} F(y, t)}{\left[\frac{\partial F}{\partial y} + q F \right]_{y=0}} dt \quad (46)$$

4. Asymptotic Form of the Solution

Various asymptotic approximations are now investigated. For large values of the real part of $(y-t + yg)$, the function $F(y, t)$ of equation (40) is given approximately by the W.K.B. solution of equation (40). This is

$$F(y, t) \cong \frac{C e^{-i\pi/4}}{[y-t + yg(y)]^{\frac{1}{4}}} \exp \left[-i \int_t^y [u-t + ug(u)]^{\frac{1}{2}} du \right] \quad (47)$$

where C is a constant and the lower limit of the integral is arbitrary.

It should be noticed that if $g = 0$

$$F(y, t) = C W_1(t-y)$$

which for large $t-y$ behaves asymptotically like

$$F(y, t) \cong \frac{C e^{-i\pi/4}}{(y-t)^{\frac{1}{4}}} \exp \left[-i \int_t^y (u-t)^{\frac{1}{2}} du \right] \quad (48)$$

in accordance with equation (28).

Equation (46) for V in the case of an inhomogeneous atmosphere can now be simplified by noting $F(y, t)$, where it occurs in the denominator, can be replaced by $W_1(t-y)$ and therefore $\left(\frac{\partial F}{\partial y} + qF \right)_{y=0} = -W_1'(t) + qW_1(t)$.

In the numerator F can be replaced by the asymptotic form given by equation (46) where y is large. Thus

$$V = -\frac{1}{2} \left(\frac{x}{i\pi} \right)^{\frac{1}{2}} \int_C \left[\frac{1}{y-t+yg(y)} \right]^{\frac{1}{4}} \frac{e^{-i\Omega}}{W_1'(t) - qW_1(t)} dt \quad (49)$$

where

$$\Omega = xt + \int_0^y \sqrt{u-t+ug(u)} du .$$

The saddle point, determined by $d\Omega/dt = 0$, is thus given by $t = -p^2$ where

$$\int_0^y \frac{du}{\sqrt{u+p^2+ug(u)}} - 2x = 0 . \quad (50)$$

Thus, following the prescription of the saddle point method

$$V \cong e^{-i\Omega_0} \left(2x \frac{\partial p}{\partial y} \right)^{\frac{1}{2}} F(-p) \quad (51)$$

where

$$F(-p) = \frac{1}{\sqrt{\pi}} \int_C e^{ipt} \frac{dt}{W_1'(t) - qW_1(t)} \quad (52)$$

and

$$\Omega_0 = -xp^2 + \int_0^y [u+p^2+ug(u)]^{\frac{1}{2}} du . \quad (53)$$

The integral $F(-p)$ has been discussed on previous occasions and numerical results are available¹⁰. For large negative values of the argument (i. e. $p > 0$) it was noted that

$$F(-p) \cong \frac{2p}{p+iq} \exp.(-ip^3/3) \quad (54)$$

and therefore

$$V \cong e^{-i\Omega_0} \left(2x \frac{\partial p}{\partial y} \right)^{\frac{1}{2}} \frac{2p}{p+iq} \exp.(-ip^3/3) \quad (55)$$

must be the geometrical optical approximation and consequently, the factor $2p/(p+iq)$ can be identified as $1+R(\alpha)$ where $R(\alpha)$ is the Fresnel reflection coefficient for a plane wave incident at an angle α on the boundary. The parameter p is then related to α by

$$p = (ka_e/2)^{\frac{1}{3}} \cos a .$$

In the case of the linear atmosphere (i. e. $g(u) = 0$)

$$\begin{aligned} \Omega_0 &= -xp^2 + \int_0^y (u + p^2)^{\frac{1}{2}} du \\ &\cong p^3/3 + 2y^{3/2}/3 \end{aligned} \quad (56)$$

for p small compared to x and y . In this same case

$$p \cong \frac{y - x^2}{2x} \quad (57)$$

and consequently

$$\left(2x \frac{\partial p}{\partial y}\right)^{\frac{1}{2}} \cong 1 .$$

When $g(u) = 0$, the formulas thus reduce to those for a homogeneous atmosphere with a modified earth radius.

In the case of the inhomogeneous atmosphere equation (50) for p can be solved approximately to yield

$$p = \frac{1}{2} \int_0^y \frac{du}{\sqrt{u + ug(u)}} - x \quad (58)$$

under the condition that x and y are large but the difference $x - \sqrt{y}$ is finite. Thus

$$\left(2x \frac{\partial p}{\partial y}\right)^{\frac{1}{2}} \cong \frac{x^{\frac{1}{2}}}{[y + yg(y)]^{\frac{1}{4}}} \cong 1 \quad (59)$$

and

$$\Omega_0 \cong p^3/3 + \int_0^y [u + ug(u)]^{\frac{1}{2}} du . \quad (60)$$

For application to an actual atmosphere, it is necessary to evaluate the integral

$$p + x = \frac{1}{2} \int_0^y \frac{du}{\sqrt{u + ug(u)}} \quad (61)$$

where

$$g = \frac{a_e}{2K_0} \left(\frac{K - K_0}{h} - K'_0 \right) .$$

When the dielectric constant varies in an arbitrary, but monotonic, way

with height, the integral can be evaluated readily by numerical methods. Some simplicity is obtained if it can be assumed that the variation is exponential in form. For example if

$$K - K_0 = (K'_0/c) (1 - e^{-ch}) \quad (62)$$

it is seen that K decays exponentially from the value K_0 at the surface ($h=0$). The constant c is of the order 0.13 for an average atmosphere and c measures of the rate of decay of the dielectric constant with increasing height.¹³ Before proceeding further in this direction, it is probably preferable to show the connection of the above results with those of geometrical optics.

5. Distance to the Horizon

The distance s to the horizon from an elevated point in the atmosphere is often used conveniently as a parameter in comparing one type of vertical refractive index profile with another. Using a method based on a stationary phase principle, it readily follows from Bremmer³ (and others) that

$$s = n_0 \int_0^h \frac{a^2}{r} \frac{dh}{\sqrt{[rn(r)]^2 - n_0 a^2}} \quad (63)$$

where $r = a + h$, $n(r)$ is the refractive which is a function of r and n_0 is the surface value [i. e. $n_0 = n(a)$].

In the case of exponential atmosphere

$$n = 1 + be^{-ch} = \sqrt{K} \quad (64)$$

so that

$$n_0 = n|_{h=0} = 1 + b = \sqrt{K_0}$$

The term under the radical can be written

$$\begin{aligned} & a^2 \left[\left(1 + be^{-ch}\right)^2 \left(1 + \frac{h}{a}\right)^2 - \left(1 + b\right)^2 \right] \\ & \cong a^2 \left[1 + 2be^{-ch} + 2\frac{h}{a} - 1 - 2b \right] \\ & \cong a^2 \left[\frac{2h}{a} + 2b \left(e^{-ch} - 1\right) \right] \end{aligned}$$

under the assumption that $b \ll 1$ and $h/a \ll 1$. To the same

approximation $\frac{n_o}{r}$ in the integrand can be replaced by $\frac{1}{a}$. Consequently, the integral becomes

$$s \cong \int_0^{h_o} \frac{dh}{\sqrt{(2h/a) + 2b(e^{-ch} - 1)}} \quad (65)$$

which is equivalent to equation (61) derived from the full theory (with $p = 0$). Now, it is convenient to break the integration into three parts, such that

$$\int_0^{h_o} = \int_0^{h'} + \int_{h'}^{h''} + \int_{h''}^{h_o=h'''} \quad (66)$$

and thus

$$s = s' + s'' + s''' .$$

In the first range ($0 \cong h \cong h'$)

$$s' \cong \int_0^{h'} \frac{dh}{\sqrt{(2h/a) + 2b\left(-ch + \frac{(ch)^2}{2}\right)}} \quad (67)$$

which is valid if h' is chosen such that the $(ch)^3/6$ term in the expansion of e^{-ch} is negligible compared to ch . This is satisfied if

$$(ch')^2/6 \ll 1$$

(For example if, $ch' = \frac{1}{2}$ then $(ch')^2/6 = 1/24$).

Writing

$$s' \cong \frac{1}{\sqrt{(2/a) - 2bc}} \int_0^{h'} \frac{dh}{\sqrt{h(1 + ah)}} \quad (68)$$

we find

$$s' = \frac{1}{\sqrt{(2/a) - 2bc}} \frac{1}{\sqrt{a}} \tanh^{-1} \sqrt{\frac{ah'}{1 + ah'}} \quad (69)$$

where

$$a = \frac{bc^2}{(2/a) - 2bc} .$$

If the effective earth's radius is defined by

$$\frac{1}{a_e} = \frac{1}{a} + \frac{K'_0}{2K_0}$$

where

$$K_0 = (1 + b)^2 \cong 1 + 2b$$

and

$$K'_0 = \left. \frac{d}{dh} \left(1 + be^{-ch} \right)^2 \right|_{h=0}$$

$$\cong -2bc$$

and thus

$$\frac{1}{a_e} = \frac{1}{a} - bc$$

so that

$$s' = \frac{1}{\sqrt{2a/a_e}} \tanh^{-1} \sqrt{\frac{ah'}{1+ah'}} \quad (70)$$

where

$$a = \frac{c}{\sqrt{2/a_e}} \left(\frac{1}{a} - \frac{1}{a_e} \right)$$

if $ah' \ll 1$

$$s' \cong \frac{1}{\sqrt{2a/a_e}} \sqrt{ah'} = \sqrt{2a_e h'} \quad (71)$$

which is the usual approximation for the distance to the horizon from a height h' in a linear atmosphere. The hyperbolic tangent formula above is actually the distance to the horizon from a height h' in an atmosphere defined by

$$n = 1 + b(1 - ch + (ch)^2/2) \quad (72)$$

for $0 \leq h \leq h'$. This is an approximation to

$$n = 1 + be^{-ch}$$

as mentioned above.

We shall turn our attention to the third portion (i. e. $h'' \leq h \leq h''' = h_0$);

$$s''' = \int_{h''}^{h_0} \frac{dh}{\sqrt{(2h/a) + 2b(e^{-ch} - 1)}} \quad (73)$$

Now the term in the radical is written in the following form

$$(2h/a - 2b) + 2be^{-ch} = (2h/a - 2b) \left[1 + \frac{2be^{-ch}}{2h/a - 2b} \right] \quad (74)$$

Letting $u = (2h/a - 2b)$ it follows that

$$\begin{aligned} \frac{1}{\left[u \left(1 + \frac{2bc^{-ch}}{u} \right) \right]^{\frac{1}{2}}} &= \frac{1}{\sqrt{u}} \sum_{n=0, 1, 2, 3, \dots} A_n \frac{(2b)^n}{u^n} e^{-chn} \\ &= \sum_n A_n (2b)^n e^{-cnba} \times \left[\frac{e^{-\frac{cna}{2}u}}{u^{n+1/2}} \right] \end{aligned} \quad (75)$$

where $A_0 = 1$, $A_1 = -\frac{1}{2}$, $A_2 = \frac{1.3}{2.4}$

$A_3 = -\frac{1.3.5}{2.4.6}$, $A_4 = \frac{1.3.5.7}{2.4.6.8}$ and so on

provided $\frac{2be^{-ch}}{u} < 1$. This is satisfied if $2b < h/a \leq h''/a$ or if $h'' > 2ba \cong 2 \times 3 \times 10^{-4} \times 6360 \cong 3.9$ km for any value of c .

Actually, this is more stringent than necessary since $e^{-ch} < 1$.

Thus

$$s''' = \sum_{n=0, 1, 2, \dots} A_n (2b)^n e^{-cnba} \int_{u''}^{u_0} \frac{e^{-\frac{cna}{2}u}}{u^{n+1/2}} du \quad (a/2) \quad (76)$$

Now consider the integral which occurs in the above expression,

$$V_m(\beta u) = \int_u^\infty \frac{e^{-\beta u}}{u^{m+1/2}} du \quad (77)$$

and integrate it by parts successively to yield the recurrence relations

$$V_m(\beta u) = \frac{1}{(m - 1/2)} \left[\frac{e^{-\beta u}}{u^{m-1/2}} - \beta V_{m-1}(\beta u) \right] \quad (78)$$

$$V_{m-1}(\beta u) = \frac{1}{(m - 3/2)} \left[\frac{e^{-\beta u}}{u^{m-3/2}} - \beta V_{m-2}(\beta u) \right]$$

----- and so on, until

$$V_2(\beta u) = \frac{1}{(1/2)} \left[\frac{e^{-\beta u}}{u^{3/2}} - \beta V_1(\beta u) \right]$$

where

$$V_1(\beta u) = \int_u^\infty \frac{e^{-\beta u}}{u^{1/2}} du = \left(\frac{\pi}{\beta} \right)^{\frac{1}{2}} \operatorname{erfc}(\sqrt{\beta u}) \quad (79)$$

where

$$\operatorname{erfc}(\sqrt{\beta u}) = \frac{2}{\sqrt{\pi}} \int_{\sqrt{\beta u}}^\infty e^{-x^2} dx \quad (80)$$

is the complement of the error integral which is extensively tabulated.¹⁴

Finally

$$s'' = \frac{a}{2} \sum_n A_n (2b)^n e^{-nbca} \left[V_n(\beta u'') - V_n(\beta u_0) \right] \quad (81)$$

where

$$\beta = \frac{cna}{2}, \quad u'' = 2 \left(\frac{h''}{a} - b \right)$$

and $u_0 = 2 \left(\frac{h_0}{a} - b \right)$.

It would appear that in most cases h' can be taken equal to h'' and chosen so that $h' = h'' = 4$ km which is a convenient upper limit for the quadratic profile and a lower limit for the expansion formula. Thus $s = s' + s''$.

An alternative is to return to the integral s'' and expand e^{-ch} to terms of third order. This leads to

$$s'' = \frac{1}{\sqrt{(2/a) - 2bc}} \int_{h'}^{h''} \frac{dh}{\sqrt{h(1 + \gamma h)(1 + \delta h)}} \quad (82)$$

where

$$\gamma = \frac{bc^2}{2 \left(\frac{1}{a} - bc \right)} \quad \text{and} \quad \delta = \frac{bc}{6 \left(\frac{1}{a} - bc \right)} .$$

This integral is expressible in terms of elliptic integrals (see Pierce¹⁴ No.535 for example).

6. Concluding Remarks

The present analysis provides a strong justification for the validity of an equivalent earth radius concept when the refractive index of the atmosphere varies in a linear manner with height in agreement with the conclusions of Bremmer³, Norton¹¹, Miller⁹ and Northover¹⁶. The effect of the non-linearity becomes important at great heights as pointed out by Millington¹⁵ and others^{10, 17} who have suggested that the theory for a linear atmosphere could be modified by simply re-locating the optical horizon point. In the present paper a complete mathematical justification is given for this step provided that the variation of the refractive index with height is monotonic.

The idea that the structure of the diffraction fields in a monotonically varying or tapered atmosphere is similar to that for a homogeneous or airless atmosphere is in apparent contradiction to Carroll's theory¹⁸ of "radio twilight". However, the model he uses is not smoothly tapered since there are discontinuities in the gradients of the refractive index. In the bilinear model, for example, the atmosphere is represented by a layer of dielectric with a constant gradient up to about 30,000 feet after which the gradient is zero. Carroll obtains his solution by matching the appropriate wave functions at the atmosphere-free space interface¹⁹. In the free space region (above 30,000 feet for the bilinear model) only outgoing waves are allowed but at the top of the dielectric graded layer the two types of wave functions which are used can be identified as up-going and downcoming waves. In the present treatment which uses a modified W.K.B. method, the solution in this region does not have this double character.

Other models used by Carroll and Ring^{19, 20} such as the inverse-square and trilinear profiles can be objected to on similar grounds. It is of interest to note that the inverse-square profile is somewhat better in that the discontinuity is in the second derivative at the bounding surface and the computed attenuation, as expected, is greater. The trilinear model has two piece-wise linear sections with interfaces at 15,000 and 90,000 feet and the mode computations show less attenuation

which is probably a consequence of the strong reflections from the lower interface.

It is the opinion of this writer that the Carroll and Ring computations are correct for the models specified. However, they cannot be regarded as valid solutions for propagation in a monotonically tapered layer. It is probable if the model were chosen so that the higher-order derivatives of the refractive index are continuous at the interfaces, the mode solutions and the modified W.K.B. solutions should approach one another. Unfortunately, the poor convergence of the resulting mode solutions will make this ultimate comparison very difficult.

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